

**71539****High-Ti Mare Basalt  
10.90 g****INTRODUCTION**

See "Rake Sample Descriptions" and "Table of Rake Samples", as well as Fig. 1.

**PETROGRAPHY AND  
MINERAL CHEMISTRY**

Warner et al. (1978) reported the petrography and mineral chemistry of 71539. During the preparation of this catalog we examined thin section 71539, 5 and found it to be a well crystallized, sub-ophitic, medium-grained (0.3-0.7mm) basalt (Fig. 2). The sample is dominated by plagioclase and pink pyroxene which is overlain by ilmenite (Fig. 2). Ilmenite phenocrysts often exceed 1mm in length and exhibit "sawtooth" margins; blocky ilmenite is a

groundmass phase. No rutile and chromite exsolution lamellae were observed in the ilmenites. No olivine or armalcolite was observed. Native Fe and troilite (both < 0.1mm) are disseminated throughout and interstitial SiO<sub>2</sub> is present.

**WHOLE-ROCK CHEMISTRY**

Murali et al. (1977) reported the whole-rock composition of 71539, 1 in a study of Apollo 17 rake samples (Table 1). 71539 is classified as a Type A Apollo 17 high-Ti basalt, based on the classification of Rhodes et al. (1976) and Warner et al. (1979). This sample contains 8.6 wt% TiO<sub>2</sub> with a MG# of 33.5. Murali et al. (1977) distinguished 71539 by its low

V, TiO<sub>2</sub>, MgO, and Cr<sub>2</sub>O<sub>3</sub> contents and suggested that it formed part of a distinct compositional group. The REE profile (Fig. 3) is LREE-depleted with the HREE exhibiting a slight depletion relative to the MREE. However, the (La/Yb)<sub>N</sub> ratio is < 1. A negative Eu anomaly is present [(Eu/Eu\*)<sub>N</sub> = 0.55).

**ISOTOPE CHEMISTRY**

Paces et al. (1991) reported whole-rock Rb-Sr and Sm-Nd data for 71539, 6 (Tables 2 and 3). In addition, these same authors reported Rb-Sr and Sm-Nd data for mineral separates from 71539, 6 (Tables 4 and 5). Paces et al. (1991) report an internal isochron age of 3.75 ± 0.07 Ga for this sample. 71539



Figure 1: Hand specimen photograph of 71539,0. Small divisions on scale are in millimeters.

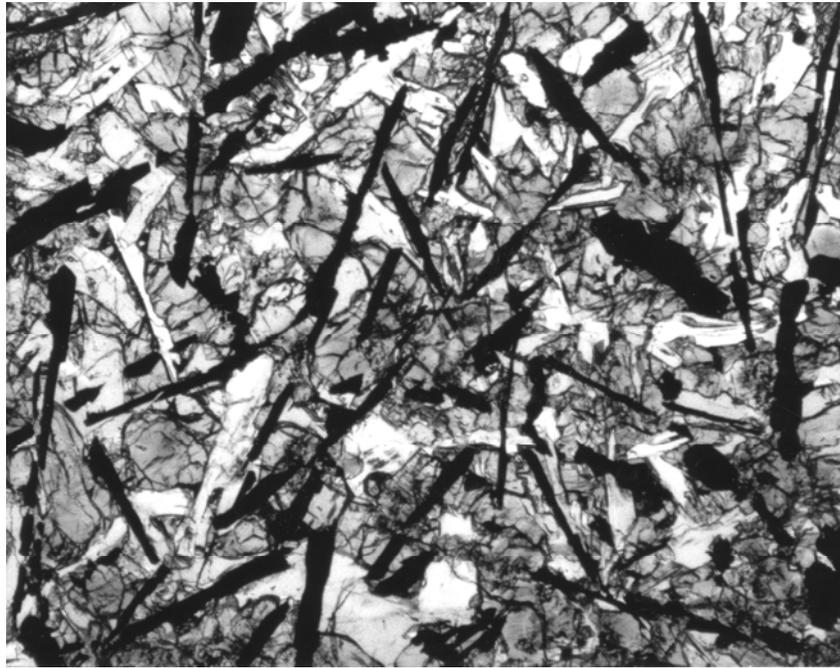


Figure 2: Photomicrograph of 71539, 5 showing ilmenite phenocrysts set in a sub-ophitic matrix. Field of view= 2.5 mm.

was studied as part of a larger isotopic investigation of the Apollo 17 high-Ti basalts.

**PROCESSING**

Of the original 10, 90g of 71539,0, a total of -10.14g remains. 71539,1 was used for INAA and the thin section ,5 was taken from this irradiated sample.

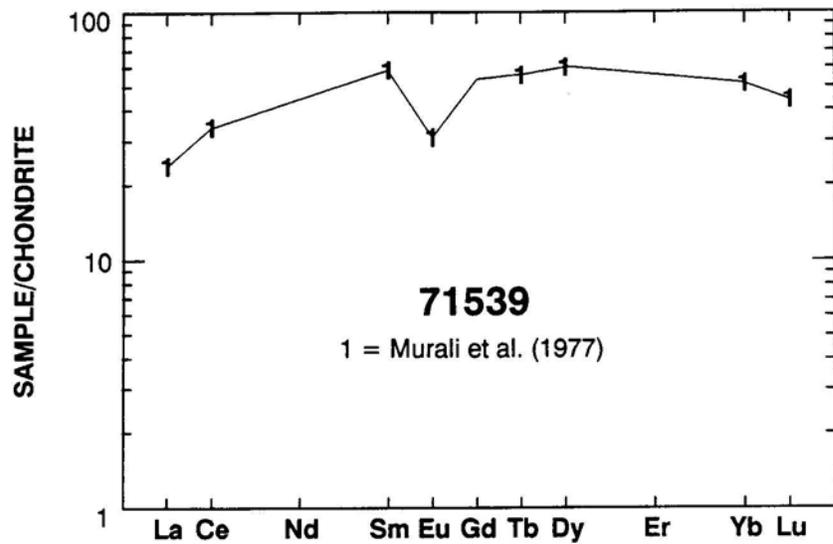


Figure 3: Chondrite-normalized rare-earth element plot of 71539. Data from Murali et al. (1977).

**Table 1: Whole-rock chemistry of 71539.**  
Data from Murali et al. (1977).

Sample 71539,1 Method N		Sample 71539,1 Method N	
SiO <sub>2</sub> (wt %)		Cu	
TiO <sub>2</sub>	8.6	Ni	
Al <sub>2</sub> O <sub>3</sub>	9.8	Co	13.5
Cr <sub>2</sub> O <sub>3</sub>	0.186	V	36
FeO	19.1	Sc	73
MnO	0.258	La	8.0
MgO	5.4	Ce	30
CaO	12.1	Nd	
Na <sub>2</sub> O	0.47	Sm	12.1
K <sub>2</sub> O	0.081	Eu	2.44
P <sub>2</sub> O <sub>5</sub>		Gd	
S		Tb	3.3
Nb (ppm)		Dy	21
Zr		Er	
Hf	9.9	Yb	11.5
Ta	1.8	Lu	1.52
U		Ga	
Th		F	
W		Cl	
Y		C	
Sr		N	
Rb		H	
Li		He	
Ba		Ge (ppb)	
Cs		Ir	
Be		Au	
Zn		Ru	
Pb		Os	

Analysis by: N = INAA.

**Table 2: Rb-Sr data for 71539,6**  
Data from Paces et al. (1991).

Rb (ppm)	0.787
Sr (ppm)	229
$^{87}\text{Rb}/^{86}\text{Sr}$	$0.009892 \pm 98$
$^{87}\text{Sr}/^{86}\text{Sr}$	$0.699776 \pm 14$
I(Sr) <sup>a</sup>	$0.699235 \pm 19$
$T_{\text{LUNI}}^{\text{b}}(\text{Ga})$	5.2

<sup>a</sup>Initial Sr isotopic ratios calculated at 3.75 Ga using  $^{87}\text{Rb}$  decay constant =  $1.42 \times 10^{-11} \text{yr}^{-1}$ .

<sup>b</sup>Model age relative to I(Sr) = LUNI = 0.69903 (Nyquist et al., 1974; Shih et al., 1986).  
 $T_{\text{LUNI}} = 1/\lambda * \ln[(^{87}\text{Sr}/^{86}\text{Sr} - 0.69903)^{87}\text{Rb}/^{86}\text{Sr} + 1]$ .

**Table 3: Sm-Nd data for 71539,6.**  
Data from Paces et al. (1991).

Sm (ppm)	13.4
Nd (ppm)	32.5
$^{147}\text{Sm}/^{144}\text{Nd}$	$0.25018 \pm 50$
$^{143}\text{Nd}/^{144}\text{Nd}$	$0.514306 \pm 11$
I(Nd) <sup>a</sup>	$0.508094 \pm 23$
$\epsilon_{\text{Nd}}^{\text{b}}$	$6.7 \pm 0.6$
$T_{\text{CHUR}}^{\text{c}}(\text{Ga})$	4.7

<sup>a</sup>Initial Nd isotopic ratios calculated at 3.75 Ga using  $^{147}\text{Sm}$  decay constant =  $6.54 \times 10^{-12} \text{yr}^{-1}$ .

<sup>b</sup>Initial  $\epsilon_{\text{Nd}}$  calculated at 3.75 Ga using present-day chondritic values of  $^{143}\text{Nd}/^{144}\text{Nd} = 0.512638$  and  $^{147}\text{Sm}/^{144}\text{Nd} = 0.1967$ .

<sup>c</sup>Model age relative to CHUR reservoir using present-day chondritic values listed above.  
 $T_{\text{CHUR}} = 1/\lambda * [((^{143}\text{Nd}/^{144}\text{Nd} - 0.512638)/(^{147}\text{Sm}/^{144}\text{Nd} - 0.1967)) + 1]$ .

**Table 4: Rb-Sr results for 71539,6 whole rock and mineral separates used for internal isochrons.**

Data from Paces et al. (1991).

	<b>Rb (ppm)</b>	<b>Sr (ppm)</b>	<b><math>^{87}\text{Rb}/^{86}\text{Sr}</math><sup>a</sup></b>	<b><math>^{87}\text{Sr}/^{86}\text{Sr}</math><sup>b</sup></b>
WR	0.787	229	$0.009893 \pm 98$	$0.699776 \pm 14$
Plg1	0.260	623	$0.001199 \pm 11$	$0.699336 \pm 22$
Plg2	0.594	536	$0.003192 \pm 31$	$0.699382 \pm 20$
NMag1 <sup>c</sup>	0.188	253	$0.002139 \pm 21$	$0.699360 \pm 21$
Px1	0.121	58.7	$0.005917 \pm 59$	$0.699564 \pm 22$
Mag1 <sup>c</sup>	0.573	103	$0.01602 \pm 32$	$0.700119 \pm 20$
Ilm1	1.23	104	$0.03395 \pm 98$	$0.701188 \pm 20$
Ilm2	0.642	128	$0.01444 \pm 98$	$0.700033 \pm 20$

<sup>a</sup>Uncertainties (corresponding to last decimal places) reported for parent/daughter ratios reflect the magnitude of the blank correction, mass spectrometer precision and corrections for the quality of spiking.

<sup>b</sup>Normalized to  $^{86}\text{Sr}/^{88}\text{Sr} = 0.1194$ . Quoted errors include 2-sigma run precision for whole rock analyses plus an additional uncertainty of 0.00001(2-sigma) reflecting corrections for fractionation and spike contributions in total-spiked mineral separates. Nd was measured as the metal ion.

<sup>c</sup>Non-pure mineral separates consisting of predominantly "nonmagnetic" plagioclase and pyroxene in NMag and "magnetic" pyroxene and ilmenite in Mag.

**Table 5: Sm-Nd results for 71539,6 whole rock and mineral separates used for internal isochrons.**

Data from Paces et al. (1991).

	Sm (ppm)	Nd (ppm)	$^{147}\text{Sm}/^{144}\text{Nd}^{\text{a}}$	$^{143}\text{Nd}/^{144}\text{Nd}^{\text{b}}$
WR	13.4	32.5	$0.25018 \pm 50$	$0.514306 \pm 11$
Plg1	2.08	5.70	$0.2204 \pm 22$	$0.513565 \pm 19$
Plg2	5.91	15.2	$0.2352 \pm 24$	$0.513899 \pm 20$
NMag1 <sup>c</sup>		9.20		$0.513868 \pm 20$
Px1	5.12	9.22	$0.3362 \pm 17$	$0.516428 \pm 18$
Mag1 <sup>c</sup>				$0.514780 \pm 16$
Ilm1	18.4	38.3	$0.2900 \pm 14$	$0.515283 \pm 22$
Ilm2		22.4		$0.514350 \pm 19$

<sup>a</sup>Uncertainties (corresponding to last decimal places) reported for parent/daughter ratios reflect the magnitude of the blank correction, mass spectrometer precision and corrections for the quality of spiking.

<sup>b</sup>Normalized to  $^{146}\text{Nd}/^{144}\text{Nd} = 0.7219$ . Quoted errors include 2-sigma run precision for whole rock analyses plus an additional uncertainty of 0.00001(2-sigma) reflecting corrections for fractionation and spike contributions in total-spiked mineral separates. Nd was measured as the metal ion.

<sup>c</sup>Non-pure mineral separates consisting of predominantly "nonmagnetic" plagioclase and pyroxene in NMag and "magnetic" pyroxene and ilmenite in Mag.